

1 **Supplementary Data 1 & 2 for 'A record of continental collision and regional sediment**
2 *flux for the Cretaceous and Paleogene core of SE Asia: Implications for early Cenozoic*
3 *palaeogeography'*

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8 **1. Methods**

9 Samples were separated using standard crushing, magnetic and heavy liquid (SPT and DIM)
10 separation techniques (e.g. Mange & Maurer 1992). For U-Pb dating of zircon, samples were
11 handpicked and analysed by LA-ICPMS using a New Wave 213 aperture imaged frequency
12 quintupled laser ablation system coupled to an Agilent 750 quadrupole-based ICP-MS. Real time
13 data were processed using the GLITTER software package (Griffin *et al.* 2008). External zircon
14 standard PLESOVIC (TIMS reference age 337.13 ± 0.37 Ma; Sláma *et al.* 2008) was used to
15 correct for instrumental mass bias. Data were filtered using standard discordance tests with a 10%
16 cutoff; discordant data that cut Concordia within error are also included. The $^{206}\text{Pb}/^{238}\text{U}$ ratio was
17 used to determine ages <1000 Ma and the $^{207}\text{Pb}/^{206}\text{Pb}$ ratio for older grains. Data were processed
18 using Isoplot (Ludwig 2003).

19 For a classification of detrital modes of sandstones we followed the Folk (1968) scheme and
20 assessed the proportions of rock constituents using the Gazzi-Dickinson method of point counting
21 (Gazzi 1966; Dickinson 1970). Thin sections were stained for potassium feldspar and point
22 counted (n = 300). Samples were chosen to be representative of the different formations and only
23 fresh unaltered samples were used. Quartz was also used as an indicator of provenance following
24 the criteria discussed by Basu *et al.* (1975) and Smyth *et al.* (2008) and references therein. An
25 assessment of heavy mineral assemblages was undertaken only on siliciclastic formations in order
26 to provide provenance data complementing other techniques. Heavy minerals mounted in Canada
27 balsam were identified by means of optical microscopy and counted (n=200) using the line-
28 counting method (Galehouse 1971; Mange & Maurer, 1992). All samples discussed are
29 summarized in Table 1 of Supplementary Data.

30 **References**

- 31
32 BASU, A., YOUNG, S.W., SUTTNER, L.J., JAMES, W.C. & MACK, G.H. 1975. Re-evaluation of the
33 use of undulatory extinction and polycrystallinity in detrital quartz for provenance
34 interpretation. *Journal of Sedimentary Research*, **45**, 873–882.
35
36 DICKINSON, W.R. 1970. Interpreting detrital modes of graywacke and arkose. *Journal of*
37 *Sedimentary Petrology*, **40**, 695–707.
38
39 FOLK, R.L. 1968. *Petrology of sedimentary rocks*. Hemphill's, Austin, Texas, 170 pp.

- 40
41 GALEHOUSE, J.S. 1971. Point counting. *In*: CARVER, R. E. (ed.) *Procedures in Sedimentary*
42 *Petrology*. Wiley Interscience, New York, 385–407.
43
44 GAZZI, P. 1966. Le arenarie del flysch sopracretaceo dell'Appennino modenese; correlazioni con il
45 flysch di Monghidoro. *Mineralogica et Petrographica Acta*, **12**, 69–97.
46
47 GRIFFIN, W.L., PEARSON, N.J., BELOUSOVA, E.A. & SAEED, A. 2008. GLITTER: data reduction
48 software for laser ablation ICP-MS. *In*: SYLVESTER, P. (ed.) *Mineralogical Association of*
49 *Canada Short Course Series Volume 40, Vancouver, B.C. Appendix 2*, 204–207.
50
51 LUDWIG, K. 2003. Isoplot 3.0. *Berkeley Geochronology Center Special Publication*, **4**.
52
53 SLÁMA, J., KOŠLER, J., CONDON, D.J., CROWLEY, J.L., GERDES, A., HANCHAR, J.M.,
54 HORSTWOOD, M.S.A., MORRIS, G.A., NASDALA, L., NORBERG, N., SCHALTEGGER, U.,
55 SCHOENE, B., TUBRETT, M.N. & WHITEHOUSE, M.J. 2008. Plešovice zircon - A new natural
56 reference material for U–Pb and Hf isotopic microanalysis. *Chemical Geology*, 1–35.
57
58 SMYTH, H., HALL, R. & NICHOLS, G.J. 2008a. Significant volcanic contribution to some Quartz-
59 rich sandstones, East Java, Indonesia. *Journal of Sedimentary Research*, **78**, 335–356.
60
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63 **2. Petrographic descriptions of West Java Sandstones**

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65 *Middle Eocene - Ciletuh Formation*

66 Analysed samples are moderately to poorly-sorted lithic-arkose to feldspathic-litharenite
67 sandstones that are dominated by grains of volcanic origin (volcanic lithic grains and plagioclase)
68 and quartz. They are relatively immature and show a wide variety of modal compositions, but all
69 plot in the magmatic arc field. Sample 31A plots in the undissected arc field and is the most
70 immature. Other samples plot in the dissected arc field indicating a minor non-volcanic
71 contribution.

72

73 *Middle Eocene - Ciemas Formation*

74 Analysed samples are moderately sorted sub-arkose and arkosic sandstones that are dominated by
75 grains of metamorphic (quartz) origin. They also contain subordinate igneous quartz, potassium
76 feldspar and metamorphic lithic fragments. Monocrystalline quartz grains are predominantly
77 strained (undulosity $>5^\circ$) indicating a metamorphic origin. Unstrained monocrystalline quartz
78 grains commonly contain strings of fluid inclusions indicating a probable plutonic source. The
79 presence of potassium feldspar and metamorphic lithic fragments indicate contributions from
80 granitic and metamorphic sources. Samples are texturally and compositionally moderately mature
81 and plot in the recycled orogenic and transitional continental fields.

82

83 *Upper Eocene – Bayah Formation*

84 Samples are moderately sorted quartz-arenite, sub-arkose and sub-litharenite sandstones that are
85 dominated by grains of metamorphic (predominantly quartz) origin. Subordinate to these are
86 magmatic quartz, chert and potassium feldspar. Monocrystalline quartz grains are predominantly
87 strained (undulosity $>5^\circ$) probably indicating a metamorphic origin (Basu *et al.* 1975). Unstrained
88 monocrystalline quartz grains commonly contain strings of fluid inclusions and are often anhedral
89 indicating a plutonic source. A significant proportion of the total quartz is polycrystalline,
90 indicating a (low grade?) metamorphic source (Basu *et al.* 1975). The presence of potassium
91 feldspar in all samples indicates a granitic source was probably important in contributing material.
92 Samples are compositionally mature, being dominated by quartz, but texturally relatively
93 immature, being typified by angular to sub-rounded grains. Modal compositions plot in the
94 recycled orogenic and craton interior fields.

95

96 *Lower Oligocene – Cikalong Formation*

97 Samples are moderately to poorly sorted quartz arenite, sub-arkose and quartzwacke sandstones
98 that are dominated by grains of metamorphic (predominantly quartz) origin. Subordinate to these
99 are chert and potassium feldspar. Monocrystalline quartz grains are predominantly strained

100 (undulosity $>5^\circ$) probably indicating a metamorphic origin. Unstrained monocrystalline quartz
101 grains appear to be plutonic. A small proportion of polycrystalline quartz is present in all samples,
102 probably indicating a (low grade) metamorphic source (Basu *et al.* 1975). The presence of
103 potassium feldspar in all samples indicates a granitic contribution. Samples are compositionally
104 very mature, being dominated by quartz, but texturally relatively immature, being typified by
105 angular to sub rounded grains. Modal compositions plot in the craton interior field.

106

107 *Lower Oligocene – Cijengkol Formation*

108 Analysed samples are moderately to poorly-sorted quartz-arenite, sub-arkose and sub-litharenite
109 sandstones that are dominated by grains of metamorphic (predominantly quartz) origin.

110 Subordinate to these are igneous quartz, potassium feldspar and metamorphic lithic fragments.

111 Monocrystalline quartz grains are predominantly strained (undulosity $>5^\circ$) probably indicating a
112 metamorphic origin (Basu *et al.* 1975). Unstrained monocrystalline quartz grains commonly
113 contain strings of fluid inclusions and are often anhedral indicating a plutonic source.

114 Polycrystalline quartz grains are present in all samples and typically have less than 5 crystal units
115 per grain, indicating a low metamorphic grade source. Potassium feldspar in some samples

116 indicates a granitic source. Samples are compositionally very mature, being dominated by quartz
117 but texturally relatively immature, being typified by angular to sub rounded grains. Modal

118 compositions plot in the recycled orogenic and craton interior fields.

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120 **References**

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122 BASU, A., YOUNG, S.W., SUTTNER, L.J., JAMES, W.C. & MACK, G.H. 1975. Re-evaluation of the
123 use of undulatory extinction and polycrystallinity in detrital quartz for provenance
124 interpretation. *Journal of Sedimentary Research*, **45**, 873–882.

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